

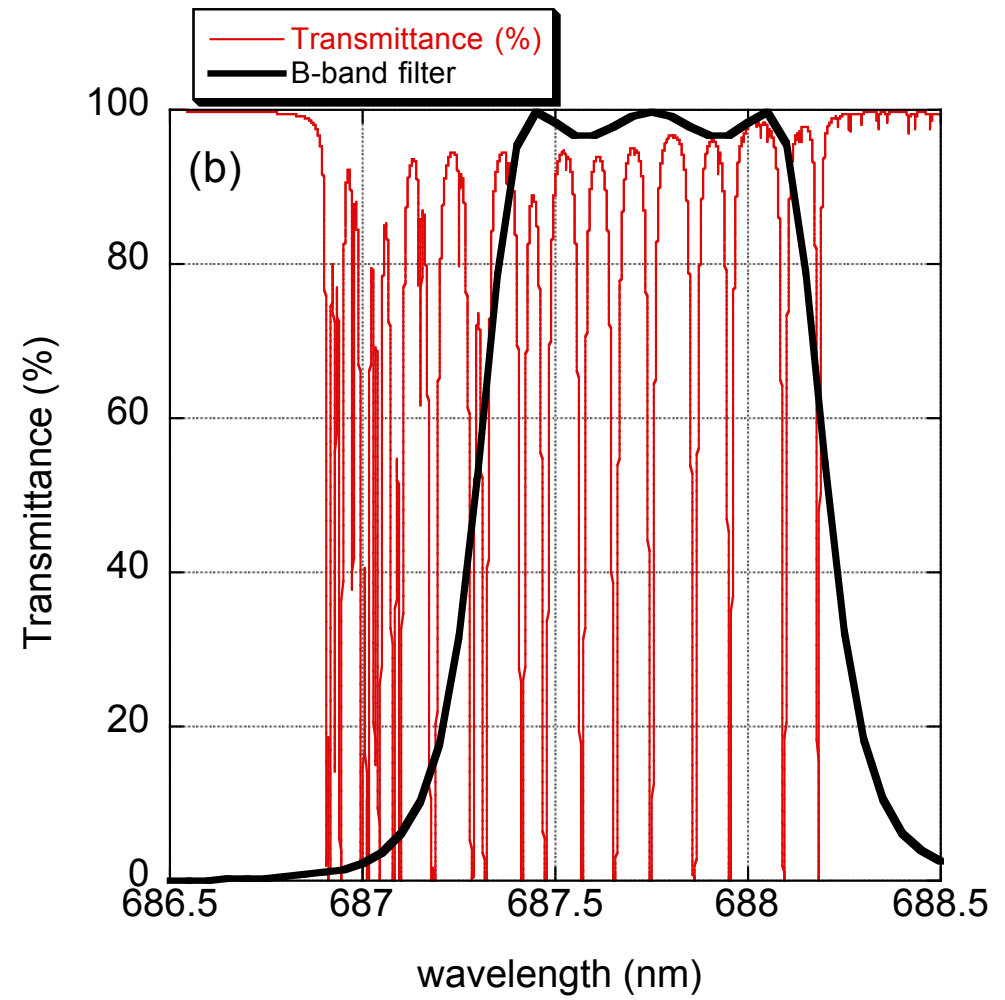
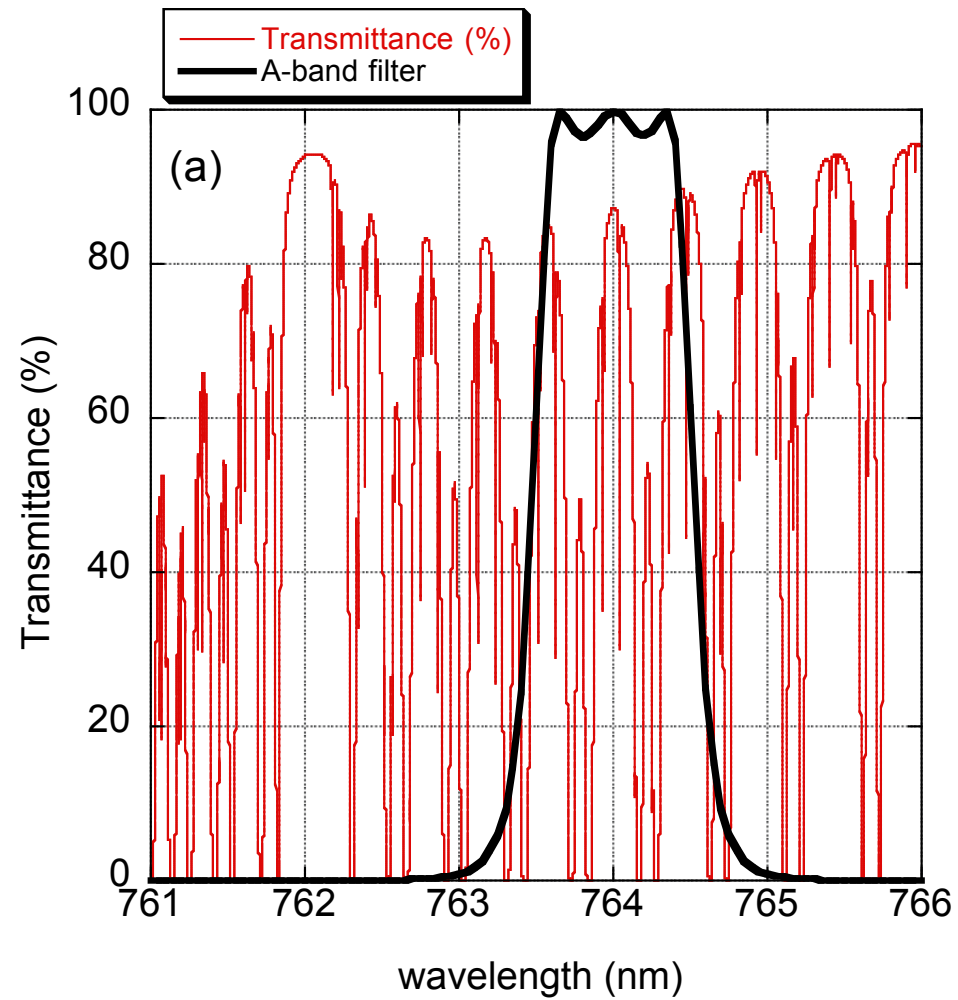
EPIC-*Inspired* Developments in Cloud and Aerosol Remote Sensing Science Using Oxygen Absorption



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Quentin Libois (CNRS/CNRM), Nicolas Ferlay (CNRS/LOA), and Alexander Marshak_(NASA – GSFC)
DSCOVR Science Team Meeting, NASA – GSFC, Greenbelt, Md, Sept. 27-29, 2022.

EPIC's take on oxygen absorption



→ Two pieces of cloud information?

Cloud centroid height from A-band: h_A

Cloud centroid height from B-band: h_B

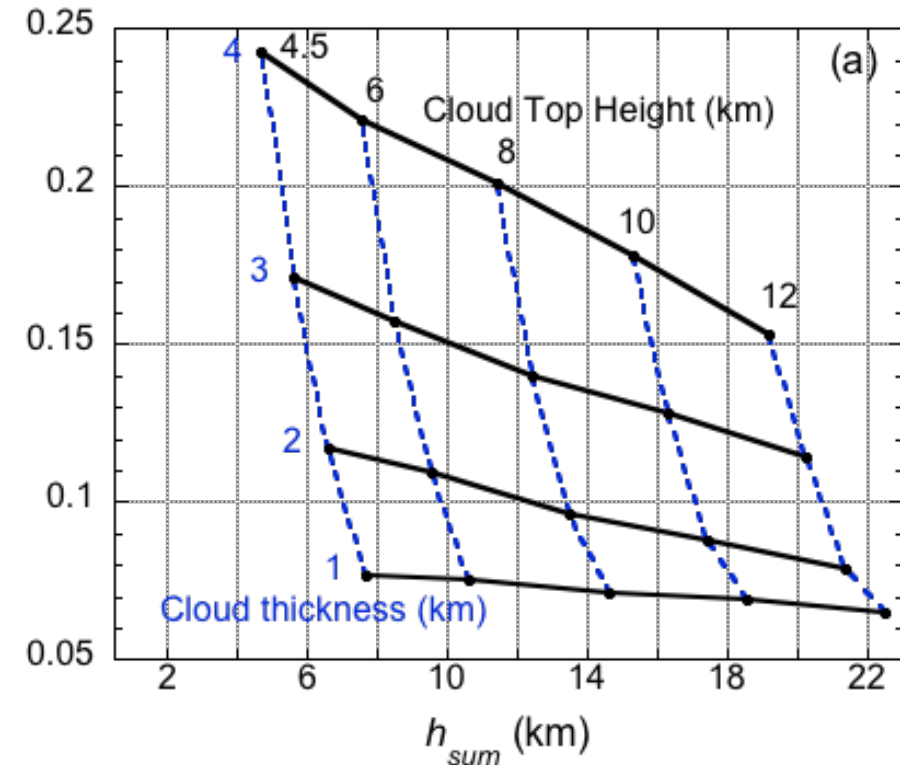
Generally,

$h_{sum} = h_A + h_B$ more sensitive to cloud top height

$h_{diff} = h_A - h_B$ more sensitive to cloud geometrical thickness

h_{diff} (km)

Yang, Y., Marshak, A., Mao, J., Lyapustin, A., and Herman, J. (2013). A Method of Retrieving Cloud Top Height and Cloud Geometrical Thickness with Oxygen A and B Bands for the Deep Space Climate Observatory (DSCOVR) Mission: Radiative Transfer Simulations. *J. Quantitative Spectrosc. Radiative Transfer* **122**, 141–149. doi:10.1016/j.jqsrt.2012.09.017



→ Two pieces of cloud information?

Cloud centroid height from A-band: h_A

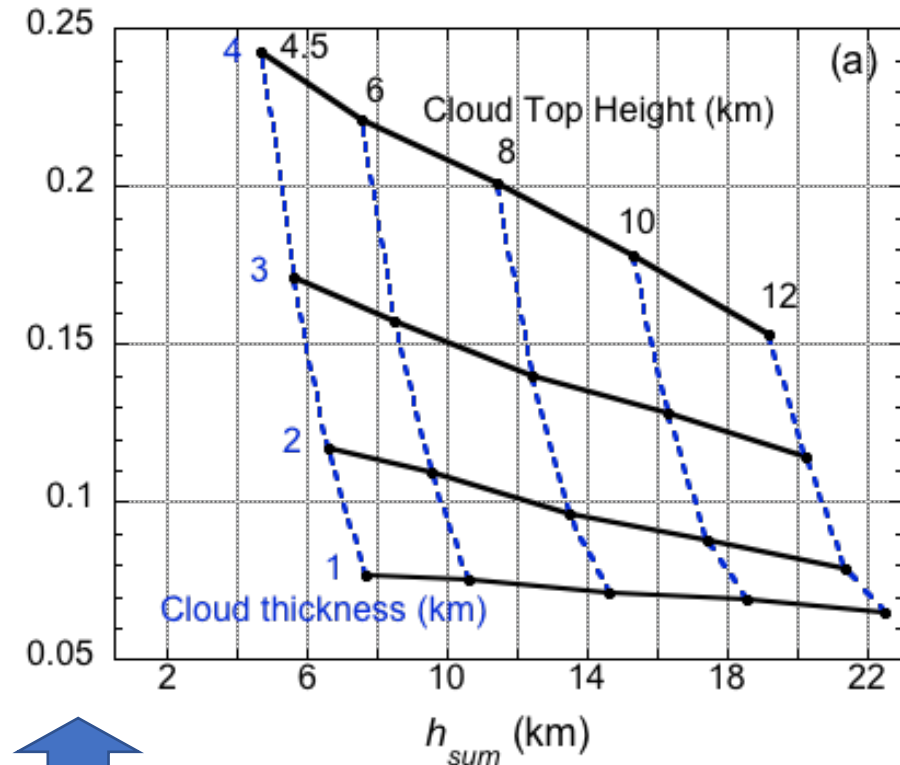
Cloud centroid height from B-band: h_B

Generally,

$h_{sum} = h_A + h_B$ more sensitive to cloud top height

$h_{diff} = h_A - h_B$ more sensitive to cloud geometrical thickness

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...unfortunately, sensor noise swamps this difference

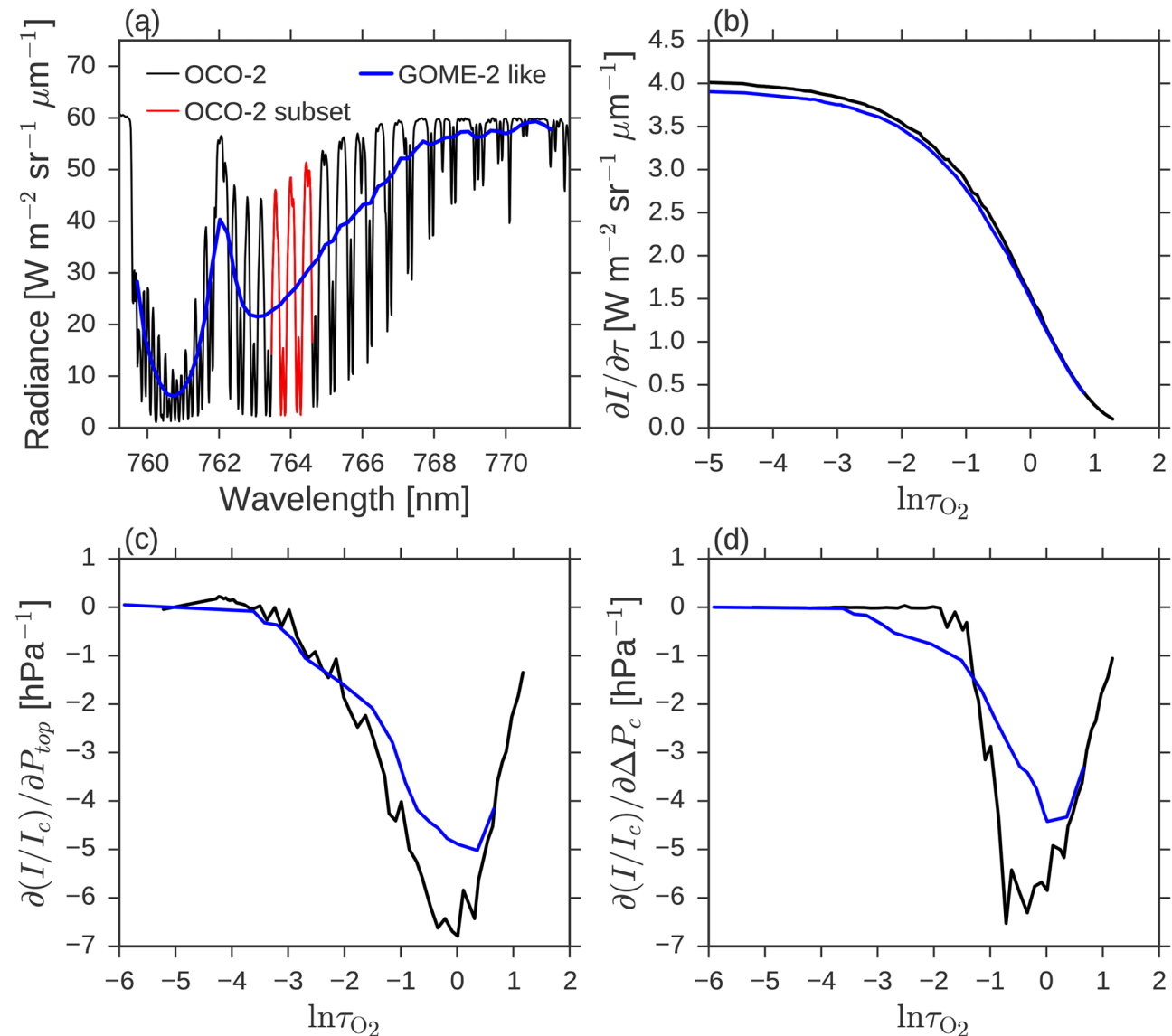
Davis, A. B., Merlin, G., Cornet, C., Labonnote, L. C., Riédi, J., Ferlay, N., et al. (2018b). Cloud Information Content in EPIC/DSCOVR's Oxygen A- and B-Band Channels: An Optimal Estimation Approach. *J. Quantitative Spectrosc. Radiative Transfer* **216**, 6–16.

doi:10.1016/j.jqsrt.2018.05.007

Davis, A. B., Ferlay, N., Libois, Q., Marshak, A., Yang, Y., and Min, Q. (2018a). Cloud Information Content in EPIC/DSCOVR's Oxygen A- and B-Band Channels: A Physics-Based Approach. *J. Quantitative Spectrosc. Radiative Transfer* **220**, 84–96. doi:10.1016/j.jqsrt.2018.09.006

GOME-2 & OCO-2/3's
 takes on the oxygen
 A-band
 (~200 to ~2000 of
 spectral channels:
 $\Delta\lambda = 0.21$ & 0.02 nm)

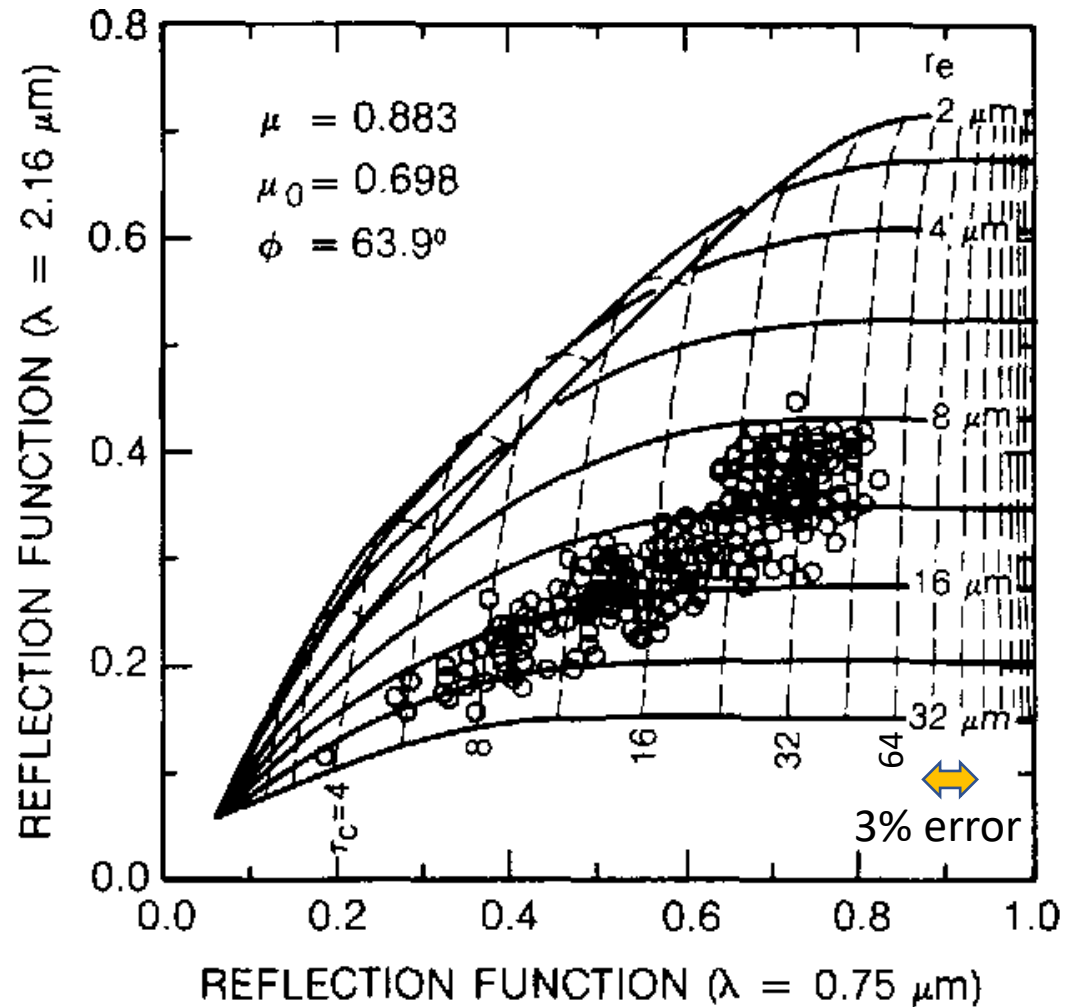
**Jacobians of $[I/I_c](\tau_{O_2})$ for cloud top
 pressure and pressure thickness \rightarrow**



Richardson, M., and Stephens, G. L. (2018). Information Content of OCO-2 Oxygen A-Band Channels for Retrieving marine Liquid Cloud Properties. *Atmos. Meas. Tech.* **11**, 1515–1528. doi:10.5194/amt-11-1515-2018

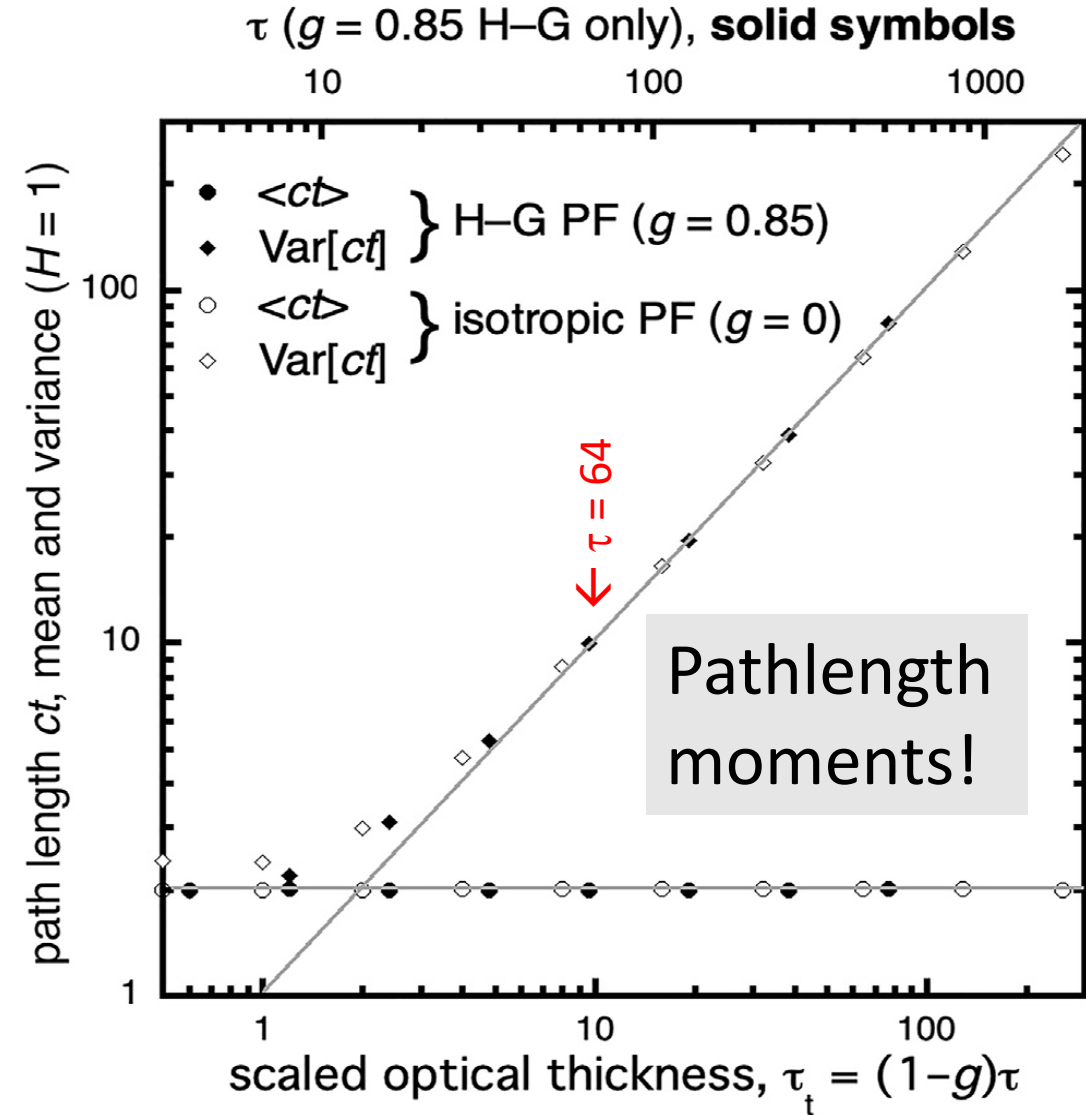
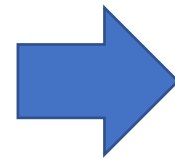
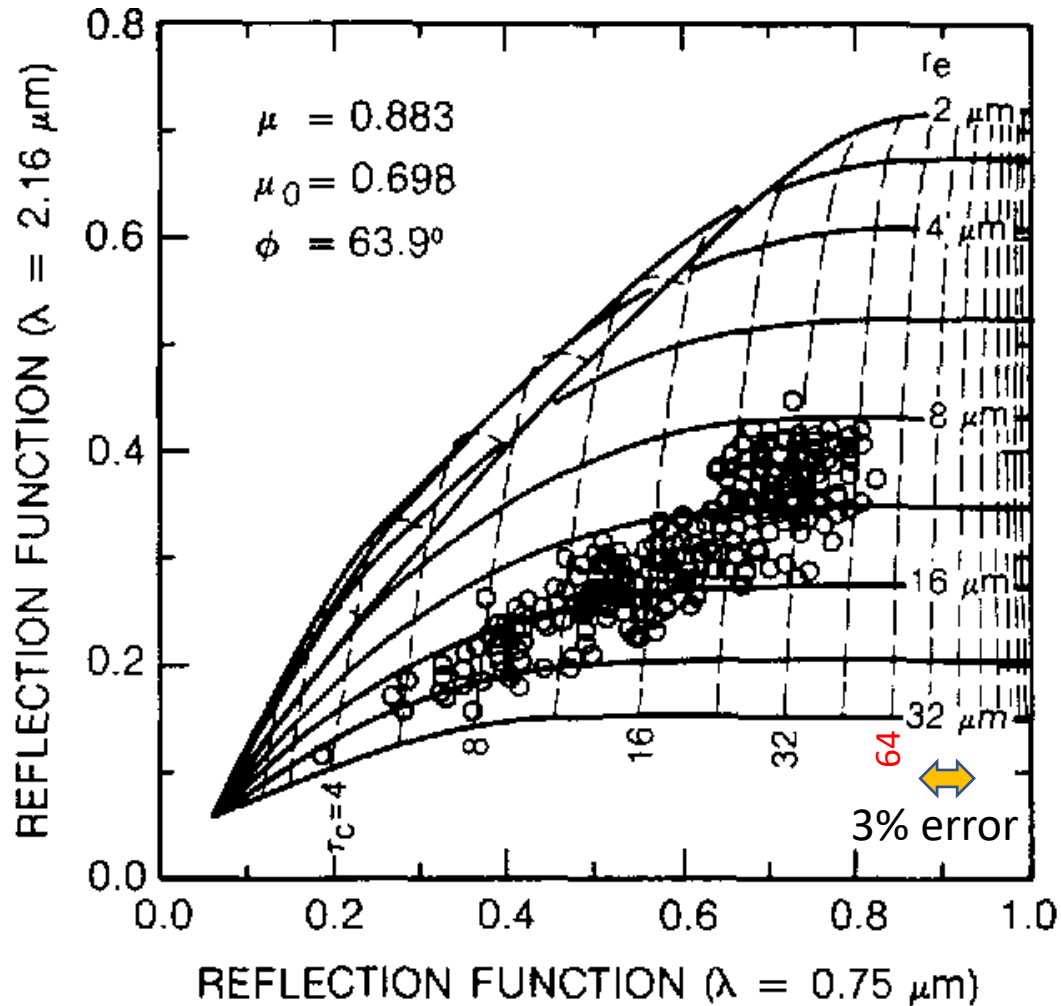
Richardson, M., Leinonen, J., Cronk, H. Q., McDuffie, J., Lebsock, M. D., and Stephens, G. L. (2019). Marine Liquid Cloud Geometric Thickness Retrieved from OCO-2's Oxygen A-Band Spectrometer. *Atmos. Meas. Techn.* **12**, 1717–1737. doi:10.5194/amt-12-1717-2019

Next up: cloud *optical* thickness from O₂ absorption



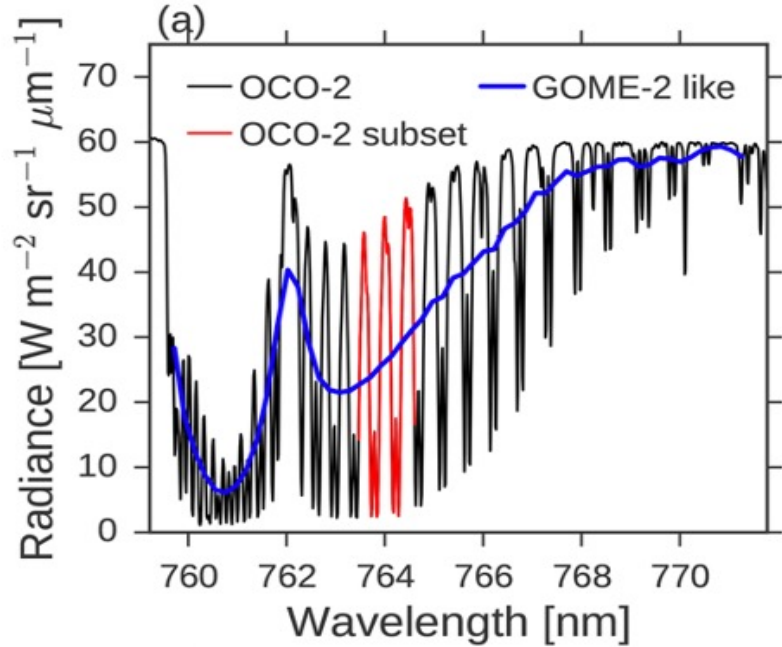
Sensitivity to COT decreases drastically as it exceeds ~50

Next up: cloud *optical* thickness from O₂ absorption



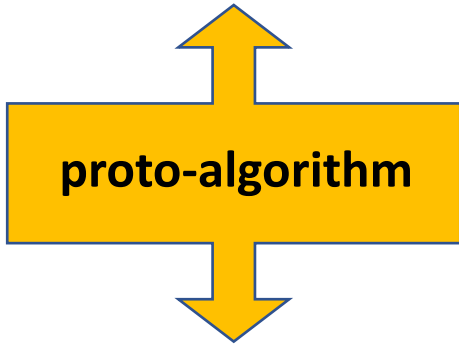
Sensitivity to COT decreases drastically as it exceeds ~50

Next up: cloud *optical* thickness from O₂ absorption



$$k_{O_2}(\lambda) = \frac{\tau_{O_2}(\lambda)}{H_{\text{mol}}}$$

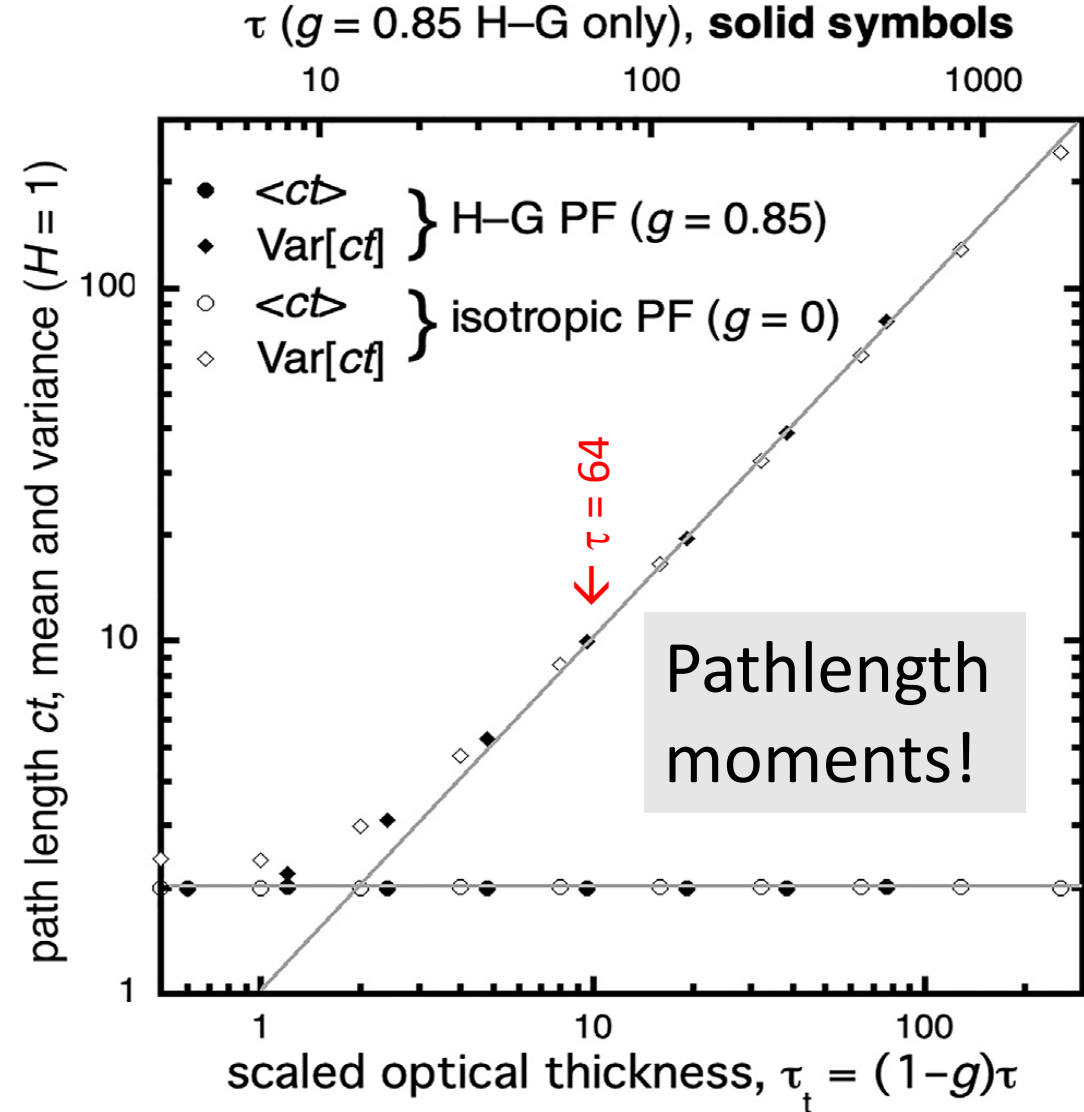
with $H_{\text{mol}} \approx 8 \text{ km}$



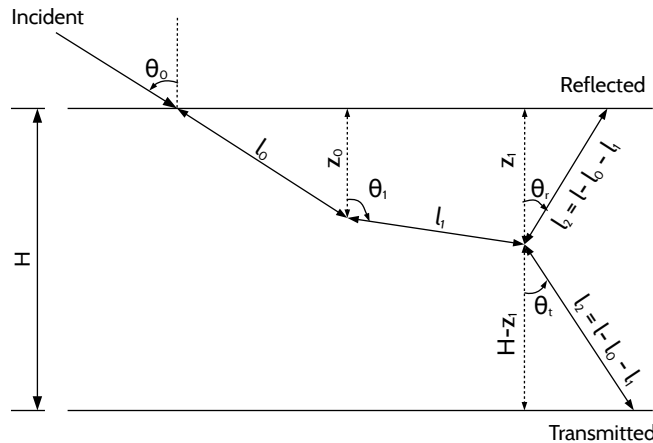
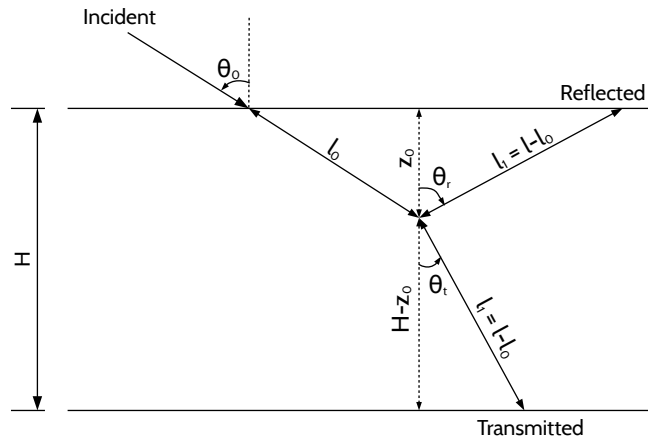
$$\ln\left(\frac{I}{I_c}\right)[k_{O_2}(\lambda)] = 0 + \left.\frac{d}{dk_{O_2}} \ln\left(\frac{I}{I_c}\right)\right|_{k_{O_2}=0} k_{O_2}(\lambda) + \frac{1}{2} \left.\left(\frac{d}{dk_{O_2}}\right)^2 \ln\left(\frac{I}{I_c}\right)\right|_{k_{O_2}=0} k_{O_2}^2(\lambda) + \dots$$

$$I(\lambda) = I_c \int_0^\infty p(ct) e^{-ctk_{O_2}(\lambda)} dct \quad \leftarrow \text{“equivalence” theorem}$$

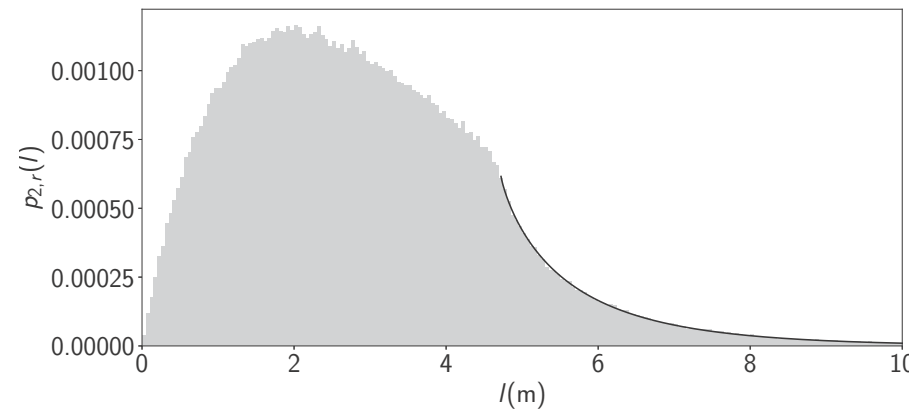
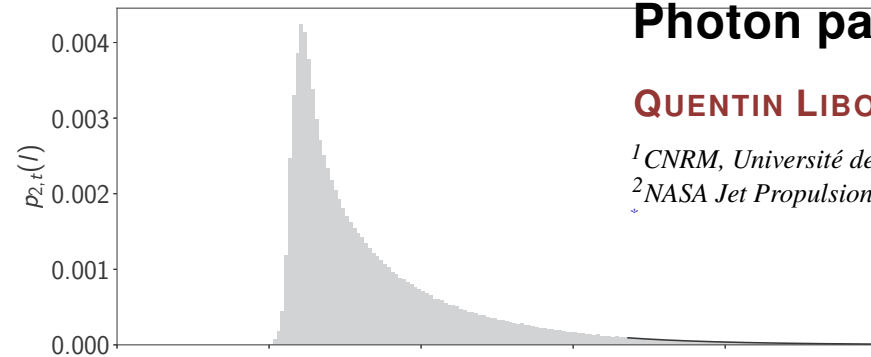
$$\Rightarrow \ln\left(\frac{I}{I_c}\right)[k_{O_2}(\lambda)] = -\langle ct \rangle k_{O_2}(\lambda) + \langle (ct)^2 - \langle ct \rangle^2 \rangle k_{O_2}^2(\lambda) + \dots$$



What about optically *thin* layers, that is, cirrus or aerosols?



$H = 2.0 \text{ m}$ $\tau = 1.0$ $\mu_0 = 0.9$ $\mu_r = \mu_t = 0.8$

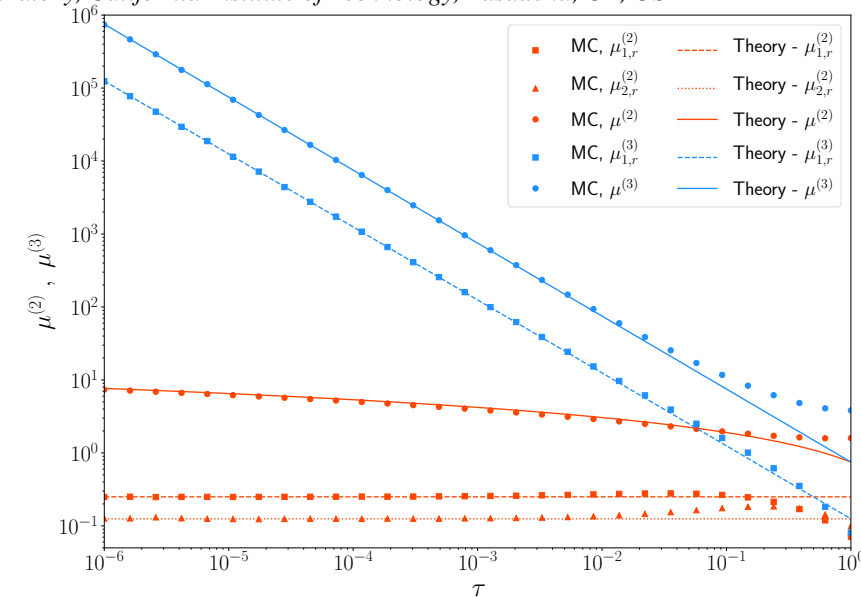


Photon path distributions in optically thin slabs

QUENTIN LIBOIS^{1,*} AND ANTHONY B. DAVIS²

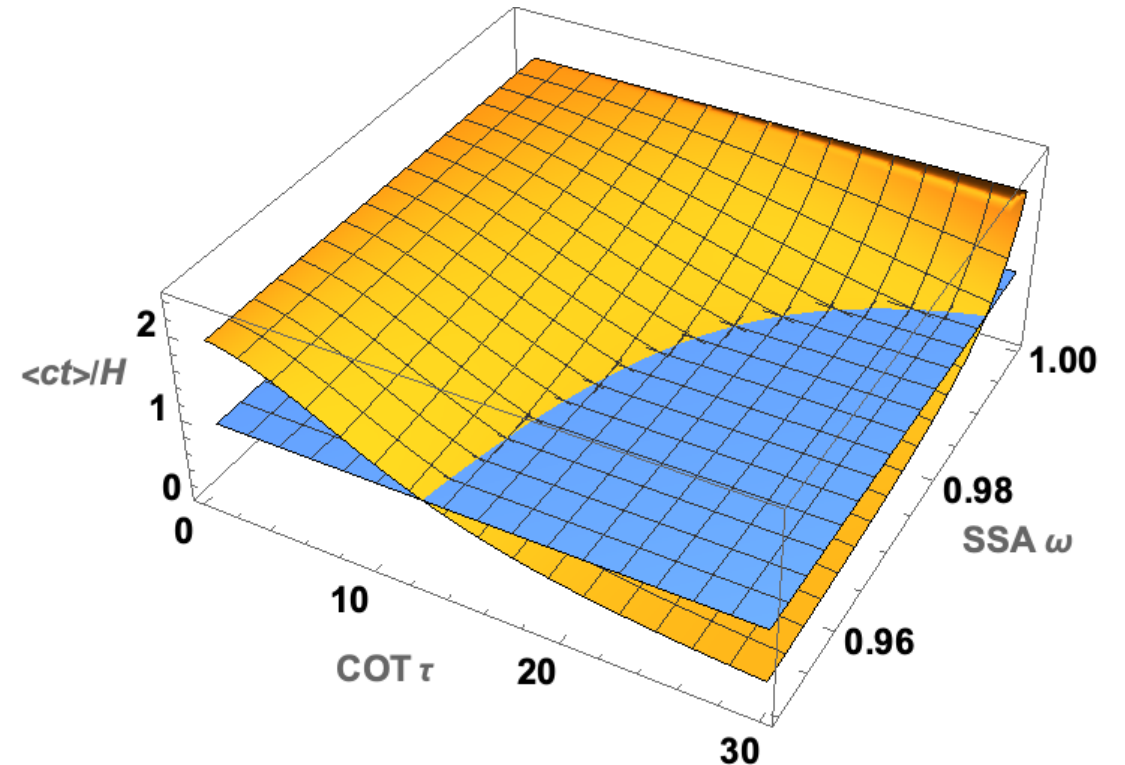
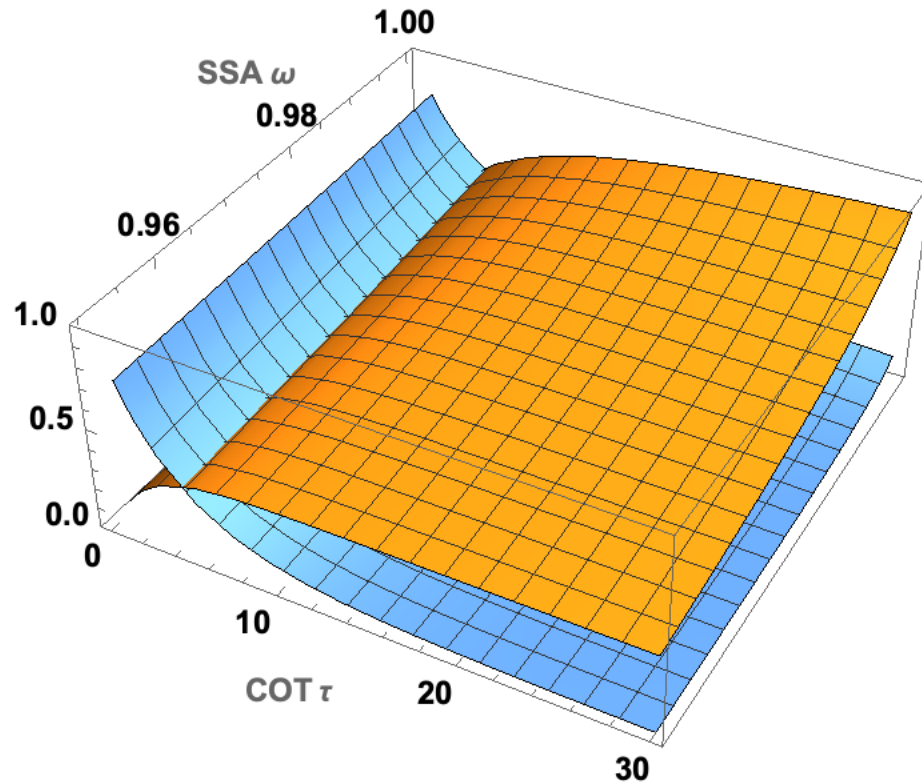
¹CNRM, Université de Toulouse, Météo-France, CNRS, Toulouse, France

²NASA Jet Propulsion Laboratory, California Institute of Technology, Pasadena, CA, USA



As for clouds, geometrical and optical thicknesses can be inferred from the mean and variance of path length. Plus we have (complicated but) analytical expressions for the distributions as well as for the moments.

What about optically *thick* absorbing aerosols plumes, say, wildfire smoke or volcanic ash?

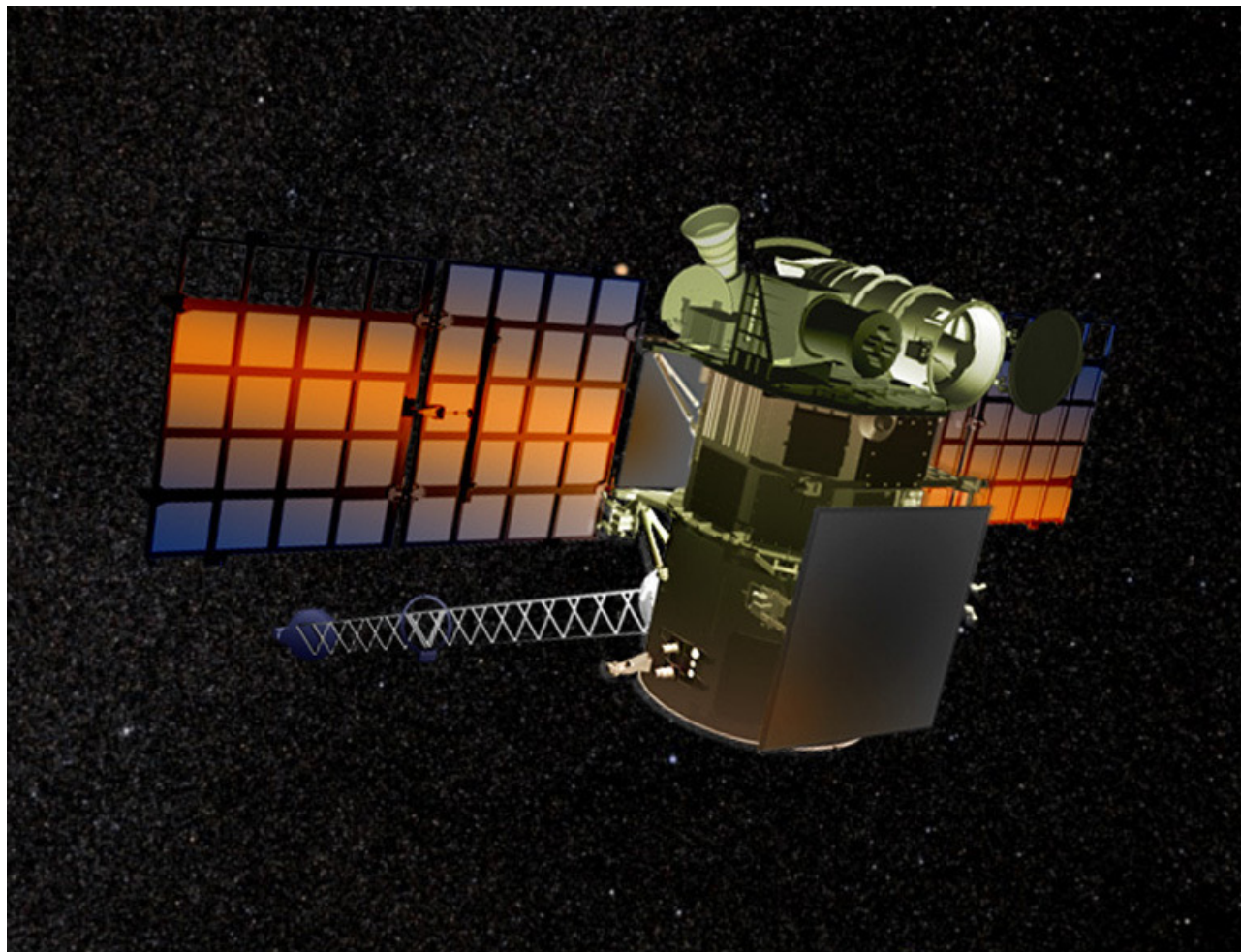


Mean pathlength $\langle ct \rangle$ is strongly affected by particulate absorption (SSA here, $\omega < 1$) in the optically thick regime.
→ For given (τ, H) , $\langle ct \rangle$ depends mainly on the SSA.
Thus, knowing $\langle ct \rangle$, puts strong constraints on the *particulate* absorption in the medium.

EPIC-inspired research on cloud and aerosol remote sensing techniques: Summary & vision

- Pathlength statistical moments for sunlight in scattering and absorbing particulate layers can be used to infer cloud-top height (CTH)/pressure, cloud geometrical/pressure thickness, and cloud optical thickness (COT). Specifically:
 - need two moments of pathlength to obtain (CTH,COT);
 - one more for CTH;
 - works for aerosols in optically thin limit as well.
- Need high-enough spectral resolution:
 - Use spatial heterodyne spectroscopic (SHS) technique to get high spectral resolving power with small imaging sensors;
 - SHS is amenable to cubesat platforms.

Thank you!



Questions?

